

EXPLANATION

ROCKS ABOVE MAIN PART OF STONY CREEK FAULT ZONE

- Q1 Alluvium, unconsolidated clay, silt, sand, and gravel, poorly sorted and poorly stratified
- Q2 Unstable debris, Arroyo indicates direction of movement
- T1 Tehama Formation, unconsolidated, blue-green claystone containing beds and lenses of poorly indurated conglomerate, sandstone and siltstone. Reported only in western part of map area
- T2 Cache Formation, white to light gray, poorly bedded, mostly unconsolidated sandstone and siltstone. Reported only in southeast corner of area
- U1 Unit 1: sandstone, light to medium-colored, gray, shaly-bedded to massive, fine to coarse grained
- U2 Unit 2: sandstone and siltstone, shaly interbedded to about equal amounts
- U3 Unit 3: sandstone and siltstone, massive to shaly, shaly bedded to laminated. Sandstone beds of unit 3 are characterized by about equal proportions of quartz and feldspar, 30 to 50 percent; rock fragments, 10 to 20 percent; mica, 1 to 2 percent. Unit contains magnesian fossils of Tertiary age near the base and Cretaceous age near the top. Includes Yucca, Palo Verde, Fanch, and Goshute Formations of Kirby (1912)
- U4 Unit 4: sandstone, light to medium-colored, gray, shaly-bedded to massive, fine to coarse grained
- U5 Unit 5: sandstone, light to medium-colored, gray, shaly-bedded to massive, fine to coarse grained
- U6 Unit 6: sandstone and siltstone, shaly interbedded to about equal amounts
- U7 Unit 7: sandstone and siltstone, shaly interbedded to about equal amounts
- U8 Unit 8: sandstone and siltstone, shaly interbedded to about equal amounts
- U9 Unit 9: sandstone and siltstone, shaly interbedded to about equal amounts
- U10 Unit 10: sandstone and siltstone, shaly interbedded to about equal amounts

STRATIGRAPHY

The Wilbur Springs quadrangle in age from Late Jurassic (Thurston) to Late Cretaceous (Compton) and, in gross aspect, is a thick succession of fine-grained hemipelagic limestones overlain by isolated magmatic units of coarse grained sandstone, conglomerate, or dioritic silt of coarse-grained beds. Commonly the coarse-grained beds are restricted in lateral extent, but a few of them have wide areal distribution. Crook (1959) ascribes such a succession of rocks to deposition in deep-sea troughs in which fine-grained hemipelagic sedimentation is continuous through time. Intermittent incursions of coarse-grained material, usually the result of turbidity currents or submarine slumps, are geologically instantaneous and form limited or tongue-shaped deposits within the fine-grained hemipelagic material. From time to time, major tectonic events in the western area of the sediment basin cause a rapid influx of coarse clastic debris which is distributed widely throughout much of the trough by northward marine currents. The coarse-grained deposits are preserved as the only distinctly mappable strata.

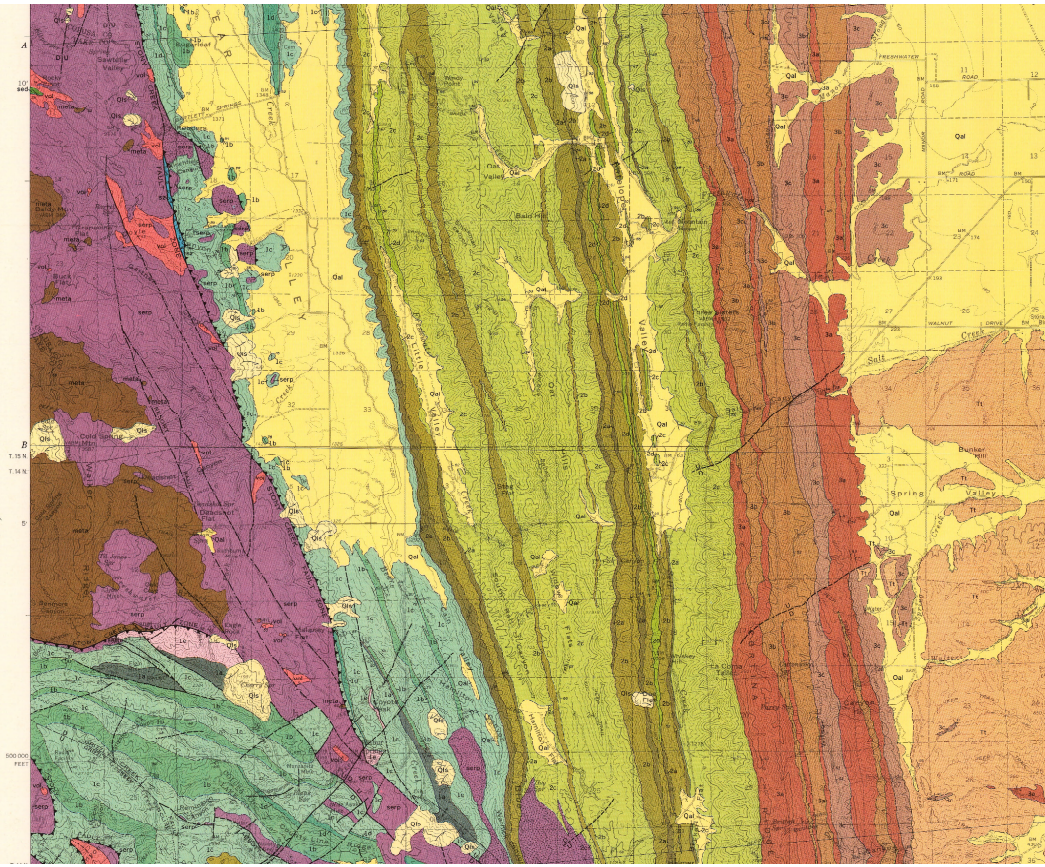
The formal stratigraphic names previously used to subdivide the thick Mesozoic sedimentary section are not adhered to in this geologic map because: (a) many of the formal names (such as Yukonite, Horsetown, Chino) for the rocks are based upon supposed age relations, rather than upon mappable lithologic boundaries; (b) several of the formal units, defined by Kirby (1912), are not mappable in the Wilbur Springs quadrangle as they pass or grade laterally into rocks of different lithology; and (c) both small- and large-scale intertonguing of units and complex facies relations exist in all parts of the section. Revisions of existing nomenclature, or definition of new stratigraphic units is beyond the scope of this publication.

The mappable unmetamorphosed strata of pre-Tertiary age are represented on the accompanying geologic map by a standardized letter symbol, A through K, denoting lithologic facies, particularly grain size. The strata are further grouped into numbered units. The units between the labeled subunits are gradational; lenses that were selected on the basis of textural criteria. The petrographic characteristics were used to define the major numbered units. The contacts between the unnumbered units are placed at the base of the lowest lithologically persistent sandstone bed that has the petrographic characteristics of the higher of the two units concerned.

The more-grained beds of each unit have similar petrologic characteristics. These characteristics, which are expressed on the relative content of quartz, plagioclase, K-feldspar, lithic fragments, and mica, are described in the map explanation. The proportion of each constituent is an average of modal point counts, restricted to 100 percent. The petrologic variations in the sandstone beds within the sequence may be generalized as follows (from base to top): (a) an increase in K-feldspar with an accompanying decrease in plagioclase to K-feldspar ratio; (b) an increase in mica; and (c) a reversible change in the percentage content of lithic fragments.

STRUCTURE

The Wilbur Springs quadrangle is divided by the Stony Creek fault into two structurally distinct portions. East of the fault zone, the Tertiary and Mesozoic succession dip homoclinally eastward under the tilted Sacramento Valley. West of the fault zone and north of the Wilbur Springs Basin, the rocks are generally horizontal. They are cutaneously metamorphosed volcanic rocks, sandstone, and siltstone. These rocks are structurally overlain by, and locally fragments of them are included in, a shaly body of sedimentite. North of the Wilbur Springs quadrangle the Stony Creek fault zone is most probably a zone of thrusts that has about the same unmetamorphosed sedimentary rocks of the Sacramento Valley stratigraphically against the metamorphic rocks (Baker and others, 1957; Pratt, 1956; Ruge, 1956; Brown, 1964; Bailey and others, 1964). The eastern trace of the fault is marked along much of its length by linear and shaly bodies of serpentinite half a mile to 2 miles wide.



omaceous mass of mass 2 are characteristic in about equal proportions of quartz and feldspar, 30-40 percent quartz, 10-15 percent feldspar, 30-40 percent mica, 10-15 percent quartz to feldspar ratio, 10-15 percent mica to feldspar ratio about 1:1. In places near the base and Commanche age near the top, include K-feldspar, Plagioclase, and Orthopyroxene of Kirby (1942).

Unit 2
 2a. sandstone, light-colored, thin-bedded to massive, fine to medium grained.
 2b. sandstone and siltstone, thin bedded to about small amounts.
 2c. sandstone and siltstone, thin bedded to massive.
 2d. conglomerate massive to thin bedded, matrix mostly composed chiefly of pebbles of chert and andesitic rocks with clasts of andesitic rocks.

Sandstone beds in lower two thirds of unit are characterized by quartz, 15-40 percent, feldspar, 10-15 percent, mica 10-15 percent, quartz to feldspar ratio, 1.5 to 2.5. Sandstone thin to upper third of unit are characterized by quartz, 30-40 percent, feldspar, 10-15 percent, mica 10-15 percent, quartz to feldspar ratio, 1.5 to 2.5, proportion to feldspar ratio about 1:1.

Contains fragments of Albian and Commanche age in upper one third of unit.

Unit 1
 1a. sandstone, pale to yellow, thin to medium bedded, fine to coarse grained.
 1b. sandstone and siltstone, thin bedded to about equal amounts.
 1c. sandstone and siltstone, dark gray to black, massive to thin bedded, sufficient near base of formation.
 1d. conglomerate, massive to thin bedded, composed chiefly of pebbles of chert and andesitic rocks with clasts of andesitic pebbles and cobbles of chert or quartz.
 1e. pebbles mostly from basalt, and andesitic rock.
 1f. thin bedded, dark gray to greenish black, thin bedded to massive, medium to coarse grained, composed chiefly of poorly sorted basaltic debris and abundant (?) quartzitic matrix, more or less calcareous.

Sandstone beds of unit 1 are characterized by about equal amounts of quartz and feldspar, 30 percent quartz, and 30 percent feldspar, 10-15 percent mica, 10-15 percent mica to feldspar ratio about 1:1. In places near the base and Commanche age near the top, include K-feldspar, Plagioclase, and Orthopyroxene of Kirby (1942).

Contains fragments of Albian and Commanche age in upper one third of unit.

Serpentinite
 Intensely altered and foliated serpentinite containing rounded clasts of serpentinized peridotite and areas of microcrystalline peridotite with irregular shapes of peridotite nodules.

ROCKS ASSOCIATED WITH STONY CREEK FAULT ZONE
 Volcanic rocks
 Finely to medium grained, dark gray to black, massive to thin bedded, extensively altered to serpentinite near Stony Creek fault zone. In place rocks blocks of Agassiz and chiefly dike-like, enclosed in serpentinite.

Serpentinite
 Serpentinized peridotite and serpentinite, microcrystalline. In places peridotite nodules as much as half an inch in diameter are found in the matrix. The serpentinite is massive, blocky with country rock commonly abundant. May include some diorite.

grain size. The strata are further grouped into numbered units. The contacts between the numbered units are gradational, hence, they were placed on the basis of textural criteria. The petrologic characteristics were used to define the major numbered units. The contacts between the major numbered units are placed at the base of the lowest lithologically persistent andesitic bed that has the petrologic characteristics of the higher of the two units concerned.

The coarse-grained beds of each unit have similar petrologic characteristics. These characteristics, which are expressed in the relative content of quartz, plagioclase, K-feldspar, lithic fragments, and mica, are described in the map explanation. The proportion of each constituent in an average modal point counts, recalculated as follows (from base to top): (a) an increase in K-feldspar with an accompanying decrease in plagioclase to K-feldspar ratio; (b) an increase in mica; and (c) a reversible change in the percentage content of lithic fragments.

STRUCTURE
 The Wilbur Springs quadrangle is divided by the Stony Creek fault zone into two structurally dissimilar terranes. East of the fault zone, the little deformed andesitic, siltstone, and conglomerate beds of the late Mesozoic succession dip homoclinally eastward and are the alluvial Sacramento Valley. West of the fault zone and south of the Wilbur Springs Basin, the rocks are strongly deformed. They are tectonically metamorphosed volcanic rocks, sandstone, and siltstone. The Franciscan assemblage of Bailey and others (1964). These rocks are structurally complex, and locally fragments of them are engulfed in a sheetlike body of serpentinite. North of the Wilbur Springs quadrangle the Stony Creek fault zone is most probably a zone of strata that has placed the unconformably sedimentary rocks of the Sacramento Valley structurally against the Franciscan rocks (Bailey and others, 1964; Page, 1966; Brown, 1964; Bailey and others, 1964). The eastern base of the three is a fault, along much of its length by linear and sheetlike bodies of serpentinite half a mile to 2 miles wide.

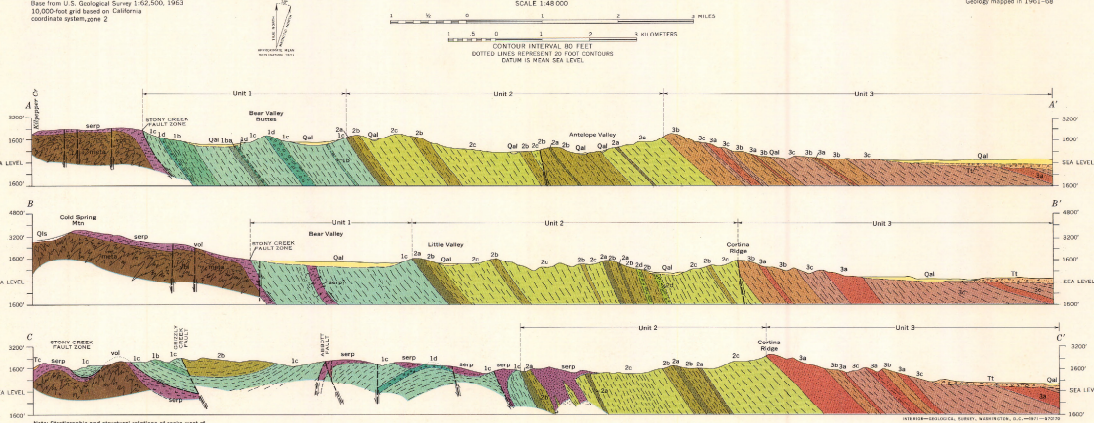
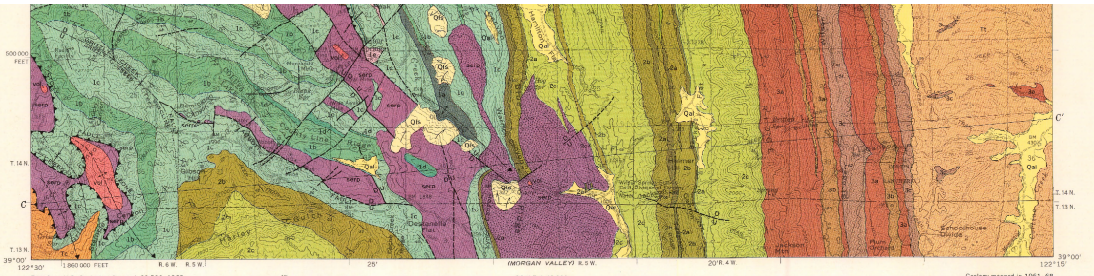
Within the Wilbur Springs quadrangle, the Stony Creek fault zone is extremely complex and is necessarily simplified on the geologic map. The surface trace parallels the western margin of Bear Valley, but south and west of the Wilbur Springs Basin it has been sharply folded, offset by the Beartown fault, and probably includes much of the southwestern part of the map.

Serpentinite has been injected into and overthrusts blocks of the late Mesozoic sedimentary succession along a branch of the Stony Creek fault zone, near 18, 19, 20, T. 13 N., R. 3 W. The contact relations are clearly exposed in Shinnelfield Canyon, in the hills on either side of a small reservoir that divides the fault zone in section 18 and 19, and in the hill south of the main fault in section 19. The rocks east of the serpentinite are steeply tilted to overthrust; those to the west are strongly deformed and locally overturned. In the NW corner, the serpentinite overlies the sedimentary rocks and merges with the large serpentinite mass west of the Stony Creek fault zone. In the eastern part of the fault, the serpentinite is capped by serpentinite, and near the junction of sections 18 and 19 serpentinite probably extends from the Stony Creek fault zone across the upturned edge of the sedimentary strata. This isolated fault zone projects above the alluvium in sections 18 and 19 are made up of steeply dipping siltstone. The fault exposed by a thin veneer of serpentinite. The fault in the southern part of Bear Valley is covered by the alluvium, although its position may be indicated by several small outcrops. The supposed location, the assumed trace of the fault, and the serpentinite now preserved as isolated patches overlying the sedimentary rocks may indicate that serpentinite once covered most of the western and southwestern part of Bear Valley.

The southwestern part of the quadrangle is probably underlain by the Stony Creek fault zone. The serpentinite-encased patches of volcanic rock exposed in the corner of small sections near Stony Creek are probably part of this folded terrain. Between these exposures of serpentinite and Bendamella Flat, the rocks of unit 1 are lithologically broken by a network of folds and faults only a few of which could be shown at the scale of the map.

The serpentinite body east of Warnick-Lynch Canyon (secs. 20, 28, and 30, T. 14 N., R. 3 W.) and secs. 1 and 2, T. 13 N., R. 3 W. is considered here to have been introduced into the sedimentary sequence either along bedding planes or along other zones of weakness that developed during the principal episode of folding and thrusting. The serpentinite is massive blocks, of which 20-40 percent is massive serpentinite peridotite blocks as much as 10 feet in diam-

CRETACEOUS
JURASSIC AND CRETACEOUS
SERPENTINE
Volcanic rocks
Serpentinite
NSHP UNKNOW



- Volcanic rocks**
Flowy argillaceous volcanic rocks, chiefly rhyolite breccia and flow breccia, extensively altered to granitic near Stony Creek fault zone. In more conchoidal blocks of argillaceous rhyolite. (Mostly common, common in serpentine facies)
- Serpentine**
Serpentinized peridotite and serpentinite, well-developed. In place granitic rhyolite as much as half as much in diameter as the original rock. Contained with quartzite rock common. May include some diagenetic or detrital serpentine of previous eras.
- Shear zone debris**
Finely divided rock debris of mixed parentage. Cherty clasts common, but contains many fragments as much as 8 inches in diameter.
- Sedimentary rocks**
Sandstone and siltstone; sandstone, dark gray to grayish gray, medium to coarse grained with variable of thin-bedded to laminated siltstone and sandstone.
- Metasedimentary rocks**
Phyllonite and somewhat foliated argillaceous rocks. Some of the rocks of the Stony Creek fault zone. Cataclastic facies most pronounced near thrust zone.
- Contact**
Long-dashed where approximately horizontal; short-dashed where undulating or irregular; dotted where questioned.
- Fault**
Long-dashed where approximately horizontal, short-dashed where undulating, dotted where questioned. In upper plate.
- Thrust fault**
Dashed where approximately horizontal, short-dashed where undulating, dotted where questioned.
- Probable fault in serpentinite**
Lined from fragments on aerial photographs and advanced of surface and size of casts rocks, dotted where questionable.
- Syncline**
Shallow curve of axial plane.
- Anticline**
Shallow curve of axial plane.
- Inclined**
Vertical.
- Overturned**
Strike and dip of beds.
- Dip**
Dip indicates top of beds, unless shown.

The southwestern part of the quadrangle is probably underlain by the Stony Creek fault zone. The serpentinite-emplaced patches of volcanic rock exposed in the cores of small anticlines near Grizzly Creek are probably part of the tilted thrust. Between these exposures of serpentinite and Detonella Flat, the rocks of unit 1 are intricately broken by a network of folds and faults only a few of which could be shown at the scale of the map.

The serpentinite body east of Warnick-Lynch Canyon (see pp. 25, 32, and 36, T. 14 N., R. 5 W., and sec. 1 and T. 13 N., R. 5 W.) is considered here to have been intruded into the sedimentary sequence along bedding planes or along other zones of weakness that developed during the pretertiary episode of folding and thrusting. The serpentinite is foliated breccia, of which 20-40 percent is massive serpentinite peridotite blocks as much as 10 feet in diameter, and the remainder is a foliated siliceous matrix of crushed serpentinite. Elastic blocks of mesosiderite and metavolcanic rocks, as well as long thin strips of fossiliferous limestone, sandstone, and siltstone, are enclosed within the foliated serpentinite. In a few places, surface creep or landsliding has extended the outcrop area of the serpentinite, and locally covered or incorporated pre-tertiary clastic serpentinite debris. This body of serpentinite has previously been described as detrital serpentinite (Taliaferro, 1943).

The serpentinite between Warnick Canyon and Deer Creek and that exposed along the adjacent fault are nearly vertical dikes that, at the surface, blend into a complex structure similar to that described by Dickinson (1966).

AGE RELATIONSHIPS UNKNOWN

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GEOLOGIC MAP OF THE WILBUR SPRINGS QUADRANGLE, COLUSA AND LAKE COUNTIES, CALIFORNIA

By
E. I. Rich
1971

California (Wilbur Springs quad). *Geol.* 1:48,000. 1971.
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